



Shallow water equations : Modeling, numerics and applications.

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1 INTRODUCTION

- Applications
- Derivation of SWE
- Properties of SWE and numerical illustration

2 NUMERICAL APPROXIMATION OF SWE

- Finite volume scheme for Homogenous SWE
- Finite volume scheme for non Homogenous SWE
- Kinetic scheme

3 SOME PHYSICAL APPLICATIONS DERIVED FROM SWE

- A sediment transport model
- An unsteady mixed flows in closed water pipes model (PFS model)

4 CONCLUDING REMARKS

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The shallow water equations (SWE), first introduced [SV] by Adhémar Jean Claude Barré de Saint-Venant are also called **Saint Venant equations**.



FIGURE: Adhémar Jean Claude Barré de Saint-Venant (1797-1886)



J-C. B. de Saint-Venant . *Théorie du mouvement non-permanent des eaux, avec application aux crues des rivières et à l'introduction des marées dans leur lit*. Compte Rendu à l'Académie des Sciences, 1871.

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FIGURE: Adhémar Jean Claude Barré de Saint-Venant (1797-1886)

The SWE are a set of *hyperbolic partial differential equations* (or parabolic if viscous shear is considered) obtained from the Euler equations (Navier-Stokes equations) under the following assumptions :

- **large-scale assumption** ($\varepsilon = \frac{H}{L}$ where H being a characteristic height and L a characteristic length of the domain)
- and hydrostatic approximation $\partial_z p = -g$.



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River

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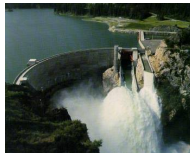
They are commonly used to describe rivers or lakes, **flooding**,



River



Flooding



Dam break

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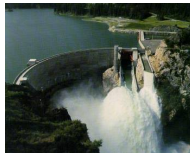
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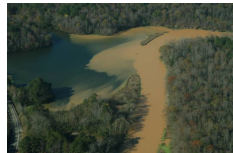
River



Flooding



Dam break



Sedimentation

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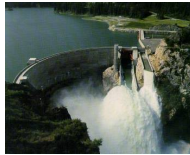
They are commonly used to describe rivers or lakes, flooding, Sedimentation/erosion flows, **channel flows**,



River



Flooding



Dam break



Sedimentation



Open channel flows



Water pipe breaking

Based on the same principle, one can find a compressible variant of Saint-Venant equation (as in [BEG11] for unsteady mixed flows in closed water pipes)



Ersoy et al., *A mathematical model for unsteady mixed flows in closed water pipes*, SCIENCE CHINA Math.,55(1),pp 1-26,2012.

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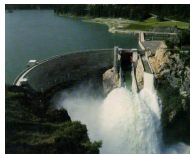
They are commonly used to describe rivers or lakes, flooding, Sedimentation/erosion flows, channel flows, **Multilayer flows**, ...



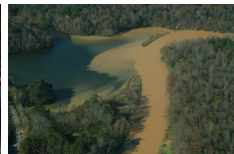
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EQUATIONS

The conservative ($S_h = S_{hu} = 0$) 1D SWE are :

$$\begin{cases} \partial_t h + \partial_x(hu) & = S_h \\ \partial_t(hu) + \partial_x(hu^2 + gh^2/2) & = S_{hu} \end{cases}$$

$h(t, x)$: density

where $u(t, x)$: velocity of the water column with S_u, S_{hu} standing for source

g : gravity strength

terms (depending on the modeling problem).

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For instance,

- S_h can be a raining term
- S_{hu} can be $-gh\partial_x Z + S_f$ where Z stands for the bathymetry (eventually depends on time) and S_f a friction term

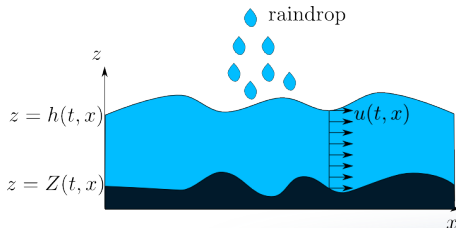


FIGURE: configuration

Almost based on the following assumptions :

- Almost flat bottom

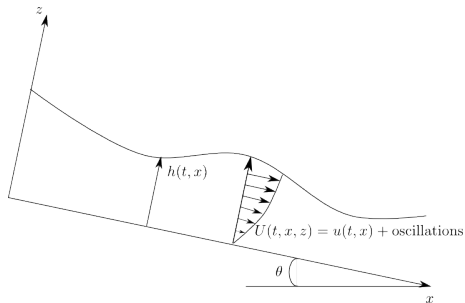


FIGURE: Almost flat bottom configuration

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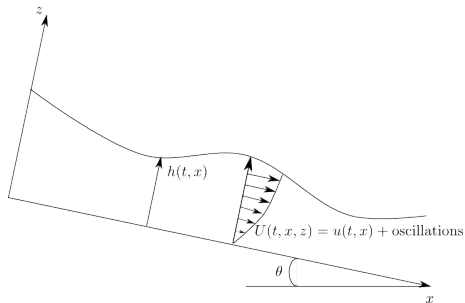


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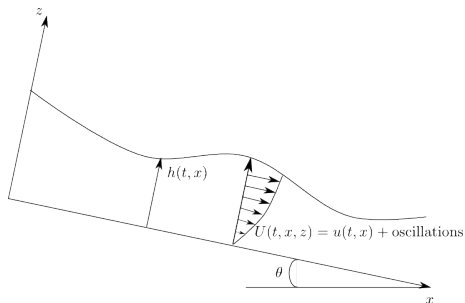


FIGURE: Almost flat bottom configuration

- Constant velocity in z
- Rectangular channel

A mathematical rigorous derivation :

- [GP01] a 1D viscous (non viscous) Saint-Venant system is obtained
- [M07] a 2D viscous (non viscous) Saint-Venant system is obtained

$$\begin{cases} \partial_t h + \operatorname{div}_{x,y}(hu) & = 0 \\ \partial_t(hu) + \operatorname{div}_{x,y}(hu \otimes u + gh^2/2I_2) & = -gh\nabla_{x,y}Z - 2\Omega \times hu - \kappa(h,u)u \end{cases}$$

where $U(t, x, y, z) = u(t, x) + O(\varepsilon)$

- ...



J-F. Gerbeau, B. Perthame. *Derivation of viscous Saint-Venant system for laminar shallow water ; numerical validation.*, Discrete Contin. Dyn. Syst. Ser. B, 1(1) :89–102, 2001.



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using

- an asymptotic expansion
- and a vertical averaging of the NS equations



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under the assumptions :

- Shallow water assumption \rightarrow hydrostatic hypothesis
- Scaled viscosity and friction \rightarrow almost constant velocity
- Almost flat bottom and free surface



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SWE

$$\begin{cases} \partial_t h + \partial_x(hu) & = 0 \\ \partial_t(hu) + \partial_x(hu^2 + gh^2/2) & = -gh\partial_x Z \end{cases}$$

- Conservation law
- Hyperbolic system (wave propagation, weak solution)
- Entropy inequality
- Positivity of water depth (invariant domain, dry zones, flooding zones)
- Non-trivial steady states

- The conservative SWE

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- The quasi-linear form of the SWE

$$\partial_t W + A(W)\partial_x W = S$$

$$\text{where } W = \begin{pmatrix} h \\ hu \end{pmatrix}, A(W) = \begin{pmatrix} 0 & 1 \\ u^2 - gh & 2u \end{pmatrix}, S = \begin{pmatrix} 0 \\ -gh\partial_x Z \end{pmatrix}$$

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- SWE system is hyperbolic $\Leftrightarrow A$ is diagonalizable on \mathbb{R} meaning that $\lambda = u \pm c \in \mathbb{R}$ where $c = \sqrt{gh}$ (sound speed).
- SW system is strictly hyperbolic on the set $\{(t, x); h(t, x) > 0\}$.

Noting the Froude number $Fr = \frac{u}{c}$, one has a

- $Fr < 1$ **Fluvial flow** (sub-critical flows)

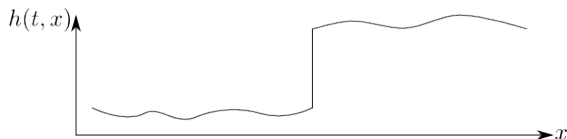
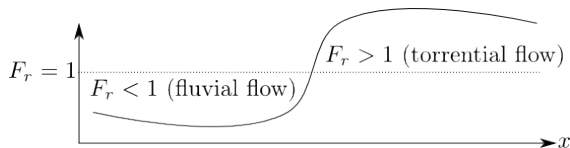
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- $Fr > 1$ **Torrential flow** (super-critical flows)

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Moreover, one has a **hydraulic bore** (discontinuous solution) whenever



a hydraulic bore (jump)



kitchen sink

It is well-known that even if the initial data are smooth, the solutions of

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Uniqueness is recovered (weak physical solution) by completing the SW system with an **entropy inequality** of the form :

$$\frac{\partial s(w)}{\partial t} + \frac{\partial \psi(w)}{\partial x} \leq 0$$

where $w = (h, u, Z)$ (non conservative variable) and (s, ψ) stands for a convex entropy-entropy flux pair

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$$s(w) = h \frac{u^2}{2} + gh \left(\frac{h}{2} + Z \right) \text{ and } \psi(w) = u(s(w) + g \frac{h^2}{2}) .$$

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- entropic scheme ,steady states preservation scheme are open problems
- **positivity of the water height** → technical difficulties.

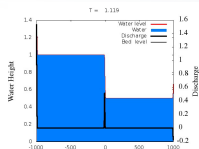


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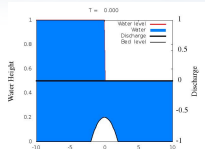


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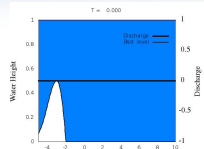
NUMERICAL ILLUSTRATION



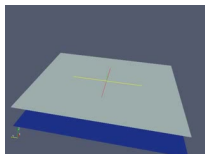
Dambreak on horizontal plane



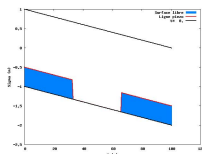
Dambreak over a non constant bottom and transcritical steady solution with shock



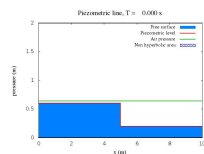
1D moving bottom



2D moving bottom



Dambreak on inclined plane and drying and flooding phenomena (closed pipe)



A air entrainment model (closed pipe)

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- Finite volume scheme for non Homogenous SWE
- Kinetic scheme

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- A sediment transport model
- An unsteady mixed flows in closed water pipes model (PFS model)

4 CONCLUDING REMARKS

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- Applications
- Derivation of SWE
- Properties of SWE and numerical illustration

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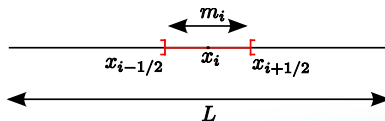


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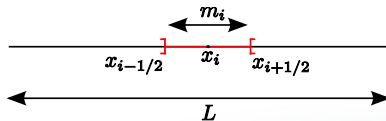


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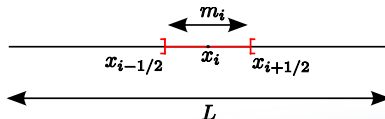


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- integration over the strip $[t_n, t_{n+1}] \times m_i$

$$u_i^{n+1} = u_i^n - \frac{\Delta t_n}{\Delta x} (F_{i+1/2} - F_{i-1/2})$$

where $t_{n+1} = t_n + \Delta t_n$ and

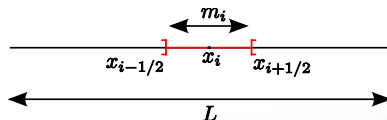


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- Explicit scheme are restricted by the time step (CFL, Courant, Friedrich and Levy condition (1928))

$$\Delta t_n \leq \frac{\Delta x}{\max_i |\lambda_i^n|}$$

- Why ?

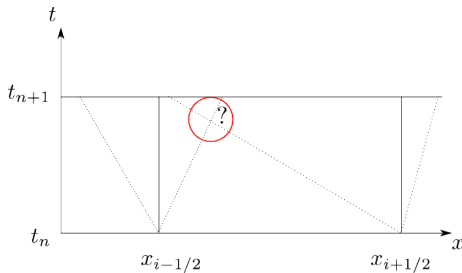


FIGURE: CFL condition

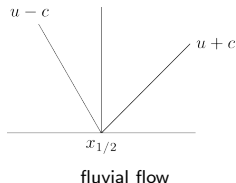
Introduction of **ghost cells** to deal with

- the most usual case
 - ▶ wall conditions ($u(x_{ghost}, t) = 0$)
 - ▶ free boundaries conditions ($u(x_{ghost}, t) = u(0, t)$ and $h(x_{ghost}, t) = h(0, t)$)

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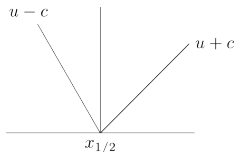
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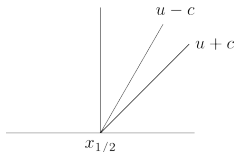
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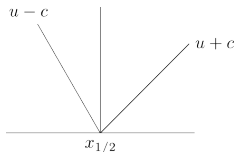


upstream torrential in

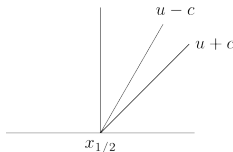
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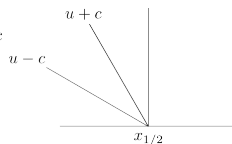
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fluvial flow



upstream torrential in



upstream torrential out

with possibly varying in time for upstream and downstream conditions.

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lead to unstable numerical scheme (even for the transport equation) !

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Solvers are

- either based on the **exact solution of the Riemann problem**, namely, the hyperbolic system with the following initial data

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- ▶ Godunov solver

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 - ▶ ...
- VFRoe (based on the **exact solution of the linearized Riemann problem**)

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- ▶ **solve locally**, for each interface $x = x_{i+1/2}$, the Riemann problem with the data

$$u^n(x) = \begin{cases} u_i^n & \text{if } x \leq x_{i+1/2} \\ u_{i+1}^n & \text{if } x > x_{i+1/2} \end{cases}$$

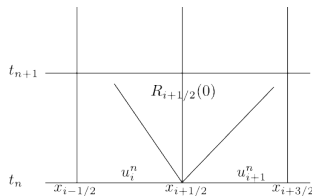


FIGURE: local Riemann problem

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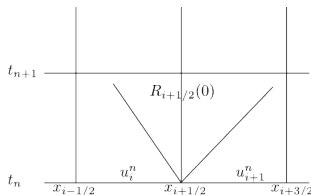


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- ▶ Call $R_{i+1/2}(x/t; u_i^n, u_{i+1}^n)$ the exact solution
- ▶ Define the numerical flux $F_{i+1/2}(u_i^n, u_{i+1}^n) := f(R_{i+1/2}(0))$

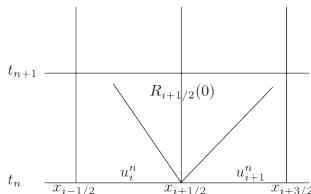


FIGURE: local Riemann problem

- **Rusanov** (also known as Lax–Friedrichs scheme)

$$F(u, v) = \frac{f(u) + f(v)}{2} - c \frac{v - u}{2}$$

for some parameter $c > 0$

- **HLL flux**

$$F(u, v) = \begin{cases} f(u) & \text{if } 0 < c_1 \\ \frac{c_2 f(u) - c_1 f(v)}{c_2 - c_1} + \frac{c_1 c_2}{c_2 - c_1} (v - u) & \text{if } c_1 < 0 < c_2 \\ f(v) & \text{if } c_2 < 0 \end{cases}$$

for some parameters $c_1 < c_2$

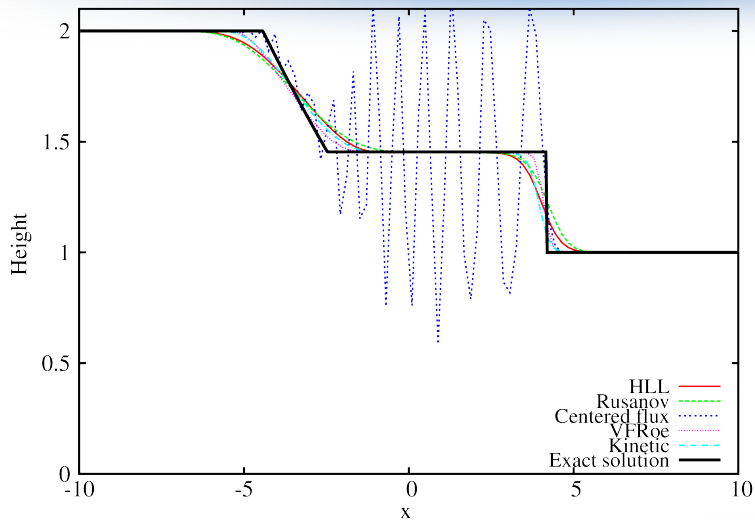


FIGURE: SWE with $u_l = 2$ and $u_r = 1$

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- Generality : **Cell-centered scheme**

$$u_i^{n+1} = u_i^n - \frac{\Delta t_n}{\Delta x} (F_{i+1/2} - F_{i-1/2}) + \Delta t_n \int_{t_n}^{t_{n+1}} \int_{m_i} S(t, x) dx dt$$

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- ▶ Consistency of the flux
- ▶ Stability properties
- ▶ **BUT usually** → bad results close to steady state solution → numerical oscillations

$$\frac{(F_{i+1/2} - F_{i-1/2})}{\Delta x} + \int_{t_n}^{t_{n+1}} \int_{m_i} S(t, x) dx dt \neq 0$$

- Generality : **Upwinded scheme** (USI methods)

$$u_i^{n+1} = u_i^n - \frac{\Delta t_n}{\Delta x} \left(F_{i+1/2}^- - F_{i-1/2}^+ \right)$$

with

$$F_{i+1/2}^-(u_i^n, u_{i+1}^n; z_{i+1} - z_i) = F_{i+1/2}(u_i^n, u_{i+1}^n) + S_{i+1/2}^-(u_i^n, u_{i+1}^n, z_{i+1} - z_i)$$

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- ▶ **Generalisation of the Lax-Wendroff theorem**

HOW TO REPRODUCE THAT ? FINITE VOLUME METHODS (NON HOMOGENOUS CASE)

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- ▶ Stability : positivity, steady states, discrete entropy inequality
- ▶ Generalisation of the Lax-Wendroff theorem
- ▶ Noting $D(u) + z = cte$ a steady state, then

THEOREM (KATSAOUNIS, PERTHAME, SIMEONI, APPLIED MATHEMATICS LETTERS, 04)

A USI scheme is well-balanced iff for all u, v, z_+, z_- such that $D(u) + z_- = D(v) + z_+$, one has

$$F(u, v) - F(u, u) + S^-(u, v, z_+ - z_-) = 0 \text{ and } F(v, v) - F(u, v) + S^+(u, v, z_+ - z_-) = 0$$

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There exist only two schemes able to :

- Hydrostatic reconstruction (Audusse, Bouchut, ...)
- **Kinetic scheme** (Perthame, Simeoni)

- Generalized Godunov schemes (Leroux, Seguin, ...)
- VFRoe schemes (Gallouët, Ersoy, ...)
- Hydrostatic reconstruction (Audusse, Bouchut, ...)
- Kinetic scheme (Perthame, Simeoni, Ersoy, ...)
- steady state profil (Greenberg, Leroux, Ersoy)
- Central upwind (Kurganov)
- ...

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Consider the SWE with a bathymetry term :

$$\begin{cases} \partial_t h + \partial_x(hu) & = & 0 \\ \partial_t(hu) + \partial_x(hu^2 + gh^2/2) & = & -ghZ'(x) \end{cases}$$

where $h(t, x)$: density
 $u(t, x)$: velocity of the water column
 $Z(x)$: topography

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 $Z(x)$: topography

THEOREM

The SW system is strictly hyperbolic for $h > 0$. It admits a mathematical entropy, which is also the physical energy $s(h, u, Z) = \frac{hu^2}{2} + \frac{gh^2}{2} + ghZ$ which satisfies the entropy inequalities :

$$\partial_t s + \partial_x \left(u \left(s + \frac{gh^2}{2} \right) \right)$$

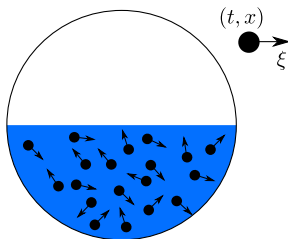
Moreover, the system admits a family of smooth steady states characterized by

$$\begin{cases} hu & = c_1 \\ \frac{u^2}{2} + g(h + Z) & = c_2 \end{cases}$$

As in gas theory ,

Describe the *macroscopic behavior* from *particle motions*, here, assumed fictitious by

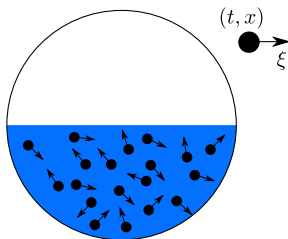
introducing $\left\{ \begin{array}{l} \text{a } \chi \text{ density function and} \\ \text{a } \mathcal{M}(t, x, \xi; \chi) \text{ maxwellian function (or a Gibbs equilibrium)} \end{array} \right.$



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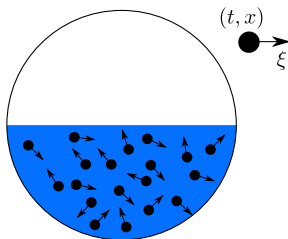
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i.e., transform the nonlinear system into a kinetic transport equation on \mathcal{M} .

Thus, to be able to define the numerical *macroscopic fluxes* from **the** microscopic one.

...Faire d'une pierre deux coups...

We introduce

$$\chi(\omega) = \chi(-\omega) \geq 0, \quad \int_{\mathbb{R}} \chi(\omega) d\omega = 1, \quad \int_{\mathbb{R}} \omega^2 \chi(\omega) d\omega = \frac{g}{2},$$

We introduce

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then we define the **Gibbs equilibrium** by

$$\mathcal{M}(t, x, \xi) = \sqrt{h(t, x)} \chi\left(\frac{\xi - u(t, x)}{\sqrt{h(t, x)}}\right)$$

PRINCIPLE

Since

$$\chi(\omega) = \chi(-\omega) \geq 0, \quad \int_{\mathbb{R}} \chi(\omega) d\omega = 1, \quad \int_{\mathbb{R}} \omega^2 \chi(\omega) d\omega = \frac{g}{2},$$

and

$$\mathcal{M}(t, x, \xi) = \sqrt{h(t, x)} \chi\left(\frac{\xi - u(t, x)}{\sqrt{h(t, x)}}\right)$$

then

MICRO-MACROSCOPIC RELATIONS

$$\begin{aligned} h &= \int_{\mathbb{R}} \mathcal{M}(t, x, \xi) d\xi \\ hu &= \int_{\mathbb{R}} \xi \mathcal{M}(t, x, \xi) d\xi \\ hu^2 + \frac{gh^2}{2} &= \int_{\mathbb{R}} \xi^2 \mathcal{M}(t, x, \xi) d\xi \end{aligned}$$

THE KINETIC FORMULATION

(h, hu) is solution of the SW system if and only if \mathcal{M} satisfy the transport equation :

$$\partial_t \mathcal{M} + \xi \cdot \partial_x \mathcal{M} - g \partial_x Z \partial_\xi \mathcal{M} = \mathcal{K}(t, x, \xi)$$

where $\mathcal{K}(t, x, \xi)$ is a collision kernel satisfying a.e. (t, x)

$$\int_{\mathbb{R}} \mathcal{K} d\xi = 0, \quad \int_{\mathbb{R}} \xi \mathcal{K} d\xi = 0.$$



B. Perthame.

Kinetic formulation of conservation laws.

Oxford University Press.

Oxford Lecture Series in Mathematics and its Applications, Vol 21, 2002.



B. Perthame and C. Simeoni

A kinetic scheme for the Saint-Venant system with a source term.

Calcolo, 38(4) :201–231, 2001.

- Recalling that Z is constant per cell

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Then $\forall (t, x) \in [t_n, t_{n+1}[\times \overset{\circ}{m}_i$

$$Z'(x) = 0$$

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- Recalling that Z is constant per cell

Then $\forall (t, x) \in [t_n, t_{n+1}[\times \overset{\circ}{m}_i$

$$Z'(x) = 0$$

\Rightarrow

$$\begin{cases} \partial_t f + \xi \cdot \partial_x f &= 0 \\ f(t_n, x, \xi) &= \mathcal{M}(t_n, x, \xi) \stackrel{def}{=} \sqrt{h(t, x)} \chi \left(\frac{\xi - u(t_n, x, \xi)}{\sqrt{h(t, x)}} \right) \end{cases}$$

by neglecting the collision kernel.

On $[t_n, t_{n+1}[\times m_i$, we have :

$$\begin{cases} \partial_t f + \xi \cdot \partial_x f & = & 0 \\ f(t_n, x, \xi) & = & \mathcal{M}_i^n(\xi) \end{cases}$$

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i.e.

$$f_i^{n+1}(\xi) = \mathcal{M}_i^n(\xi) + \xi \frac{\Delta t^n}{\Delta x} \left(\mathcal{M}_{i+\frac{1}{2}}^-(\xi) - \mathcal{M}_{i-\frac{1}{2}}^+(\xi) \right)$$

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where

$$\mathbf{u}_i^{n+1} = \begin{pmatrix} A_i^{n+1} \\ Q_i^{n+1} \end{pmatrix} \stackrel{\text{def}}{=} \int_{\mathbb{R}} \begin{pmatrix} 1 \\ \xi \end{pmatrix} f_i^{n+1}(\xi) d\xi$$

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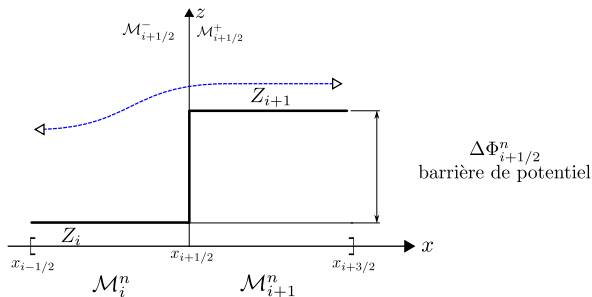
or

$$\mathbf{U}_i^{n+1} = \mathbf{U}_i^n - \frac{\Delta t^n}{\Delta x} \left(\mathbf{F}_{i+1/2}^- - \mathbf{F}_{i-1/2}^+ \right)$$

with

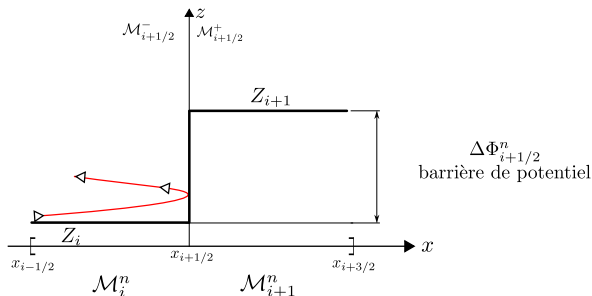
$$\mathbf{F}_{i\pm\frac{1}{2}}^\pm = \int_{\mathbb{R}} \xi \begin{pmatrix} 1 \\ \xi \end{pmatrix} \mathcal{M}_{i\pm\frac{1}{2}}^\pm(\xi) d\xi.$$

$$\mathcal{M}_{i+1/2}^{-}(\xi) = \overbrace{\mathbb{1}_{\{\xi>0\}} \mathcal{M}_i^n(\xi)}^{\text{positive transmission}} \\ + \underbrace{\mathbb{1}_{\{\xi<0, \xi^2 - 2g\Delta\Phi_{i+1/2}^n > 0\}} \mathcal{M}_{i+1}^n \left(-\sqrt{\xi^2 - 2g\Delta\Phi_{i+1/2}^n} \right)}_{\text{negative transmission}}$$



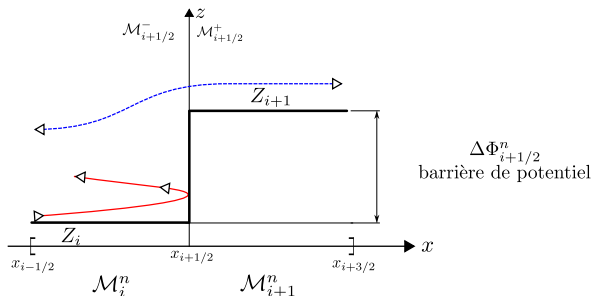
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Let us recall that we have to define a χ function such that :

$$\chi(\omega) = \chi(-\omega) \geq 0, \quad \int_{\mathbb{R}} \chi(\omega) d\omega = 1, \quad \int_{\mathbb{R}} \omega^2 \chi(\omega) d\omega = \frac{g}{2},$$

and $\mathcal{M} = \sqrt{h} \chi\left(\frac{\xi - u}{\sqrt{h}}\right)$ satisfies the equation :

$$\partial_t \mathcal{M} + \xi \cdot \partial_x \mathcal{M} - g Z' \partial_\xi \mathcal{M} = 0$$

and

$\chi \longrightarrow$ definition of the macroscopic fluxes.

One has

- **Conservativity of h** holds for every χ .
- **Positivity of A** holds for every χ **but** for numerical purpose only if $\text{supp}\chi$ is **compact** to get a **CFL** condition.

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- **discrete entropy inequalities**

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strongly depend on the choice of the χ function.

In the following, we only focus on discrete equilibrium.

STRATEGY

Set χ such that $\mathcal{M}(t, x, \xi; \chi)$ is the still water steady state solution of :

$$\xi \cdot \partial_x \mathcal{M} - g Z' \partial_\xi \mathcal{M} = 0.$$

Then

$$w\chi(w) + \{2g - w^2\} \chi'(w) = 0$$

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As a consequence, this function is the only possible choice such that

$M(t, x, \xi) = \sqrt{h}\chi\left(\frac{\xi - u}{\sqrt{h}}\right)$ satisfies the equation

$$\xi \cdot \partial_x \mathcal{M} - gZ' \partial_\xi \mathcal{M} = 0.$$

on any still water steady states,

$$u(t, x) = 0, \quad h(t, x) + Z(x) = H, \quad \forall t \geq 0.$$

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on any still water steady states,

$$u(t, x) = 0, \quad h(t, x) + Z(x) = H, \quad \forall t \geq 0.$$

Moreover, one can interpret, such a solution, as the minimum of the kinetic energy

$$E(f) = \int_{\mathbb{R}} \left(\frac{\xi^2}{2} f(\xi) + \frac{\pi^2 g^2}{6} f^3(\xi) + gZ f(\xi) \right) d\xi$$

With this χ function, one has

THEOREM

Under the CFL condition

$$\Delta t_n \max \left(|u_i^n| + \sqrt{2gh_i^n} \right) \leq \Delta x ,$$

the following assertions hold :

- ❶ *Positivity of h ,*
- ❷ *Conservativity of h ,*
- ❸ *Discrete still water steady states are preserved,*
- ❹ *Discrete in-entropy inequalities are satisfied,*
- ❺ *Drying and flooding phenomenon are naturally obtained*

As a conclusion, EVEN IF THE INTEGRAL COMPUTATION ARE NOT EXPLICIT, this kinetic scheme have all the necessary properties to get consistency, stability, convergence.

With this χ function, one has

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Under the CFL condition

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- ❺ *Drying and flooding phenomenon are naturally obtained*

In practice, I prefer to use the following simplest χ -function

$$\chi(\omega) = \frac{1}{2\sqrt{3}} \mathbb{1}_{[-\sqrt{3}, \sqrt{3}]}(\omega)$$

with

- integral computation are explicit and easy
- loosing the property 3 and a priori 4 (even if the numerical results provides good agreements with test cases)

1 INTRODUCTION

- Applications
- Derivation of SWE
- Properties of SWE and numerical illustration

2 NUMERICAL APPROXIMATION OF SWE

- Finite volume scheme for Homogenous SWE
- Finite volume scheme for non Homogenous SWE
- Kinetic scheme

3 SOME PHYSICAL APPLICATIONS DERIVED FROM SWE

- A sediment transport model
- An unsteady mixed flows in closed water pipes model (PFS model)

4 CONCLUDING REMARKS

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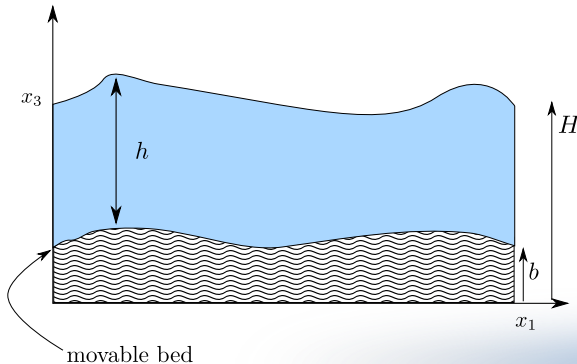
Saint-Venant equations for the hydrodynamic part :

$$\begin{cases} \partial_t h + \operatorname{div}(q) = 0, \\ \partial_t q + \operatorname{div}\left(\frac{q \otimes q}{h}\right) + \nabla\left(g \frac{h^2}{2}\right) = -gh \nabla b \end{cases} \quad (1)$$

+

a bedload transport equation for the morphodynamic part :

$$\partial_t b + \xi \operatorname{div}(q_b(h, q)) = 0 \quad (2)$$



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with

- h : water height,
- $q = hu$: water discharge,
- q_b : sediment discharge (empirical law : [MPM48], [G81]),
- $\xi = 1/(1 - \psi)$: porosity coefficient.



M. Ersoy,

Modélisation, analyse mathématique et numérique de divers écoulements compressibles ou incompressibles en couche mince,
Ph.D University of Savoie (France), 2010.



E. Meyer-Peter and R. Müller,

Formula for bed-load transport,
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A.J. Grass,

Sediment transport by waves and currents,
SERC London Cent. Mar. Technol. Report No. FL29, 1981.

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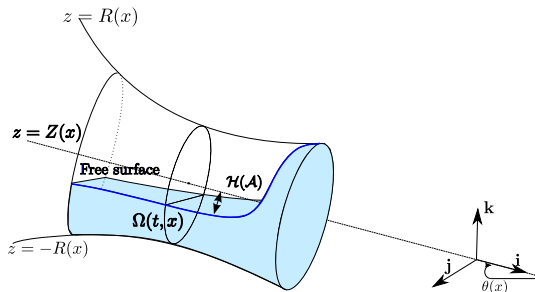
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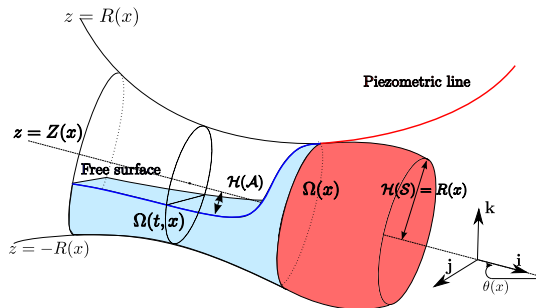
- Free surface area (SL)

sections are not completely filled and the flow is **incompressible**...



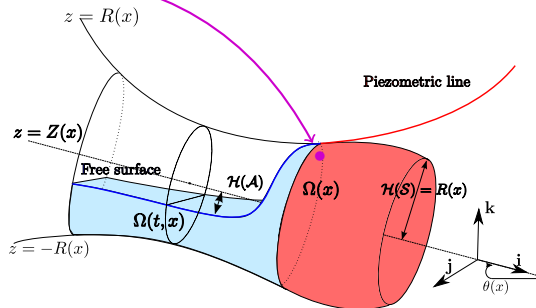
PHYSICAL APPLICATIONS : UNSTEADY MIXED FLOWS IN CLOSED WATER PIPES ?

- Free surface area (SL)
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PHYSICAL APPLICATIONS : UNSTEADY MIXED FLOWS IN CLOSED WATER PIPES ?

- Free surface area (SL)
sections are not completely filled and the flow is incompressible...
- Pressurized area (CH)
sections are non completely filled and the flow is compressible...
- Transition point



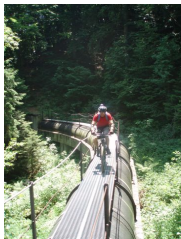
EXAMPLES OF PIPES



Orange-Fish tunnel



Sewers ...in Paris



Forced pipe



problems ...at Minnesota

<http://www.sewerhistory.org/grfx/misc/disaster.htm>

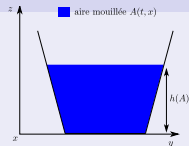
GENERALLY

Saint-Venant equations :

$$\begin{cases} \partial_t A + \partial_x Q = 0, \\ \partial_t Q + \partial_x \left(\frac{Q^2}{A} + g I_1(A) \right) = 0 \end{cases}$$

with

$A(t, x)$:	wet area
$Q(t, x)$:	discharge
$I_1(A)$:	hydrostatic pressure
g	:	gravity



Advantage

- Conservative formulation → Easy numerical implementation



Hamam and McCorquodale (82), Trieu Dong (91), Musandji Fuamba (02), Vasconcelos *et al* (06)

GENERALLY

Allievi equations :

$$\begin{cases} \partial_t p + \frac{c^2}{gS} \partial_x Q = 0, \\ \partial_t Q + gS \partial_x p = 0 \end{cases}$$

with

$p(t, x)$:	pressure
$Q(t, x)$:	discharge
$c(t, x)$:	sound speed
$S(x)$:	section

Advantage

- Compressibility of water is taking into account \Rightarrow Sub-atmospheric flows and over-pressurized flows are well computed

Drawback

- Non conservative formulation \Rightarrow Cannot be, at least easily, coupled to Saint-Venant equations

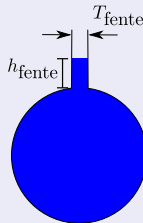


Winckler (93), Blommaert (00)

GENERALLY

Saint-Venant with Preissmann slot artifact :

$$\begin{cases} \partial_t A + \partial_x Q = 0, \\ \partial_t Q + \partial_x \left(\frac{Q^2}{A} + g I_1(A) \right) = 0 \end{cases}$$



Advantage

- Only one model for two types of flows.

Drawbacks

- Incompressible Fluid \Rightarrow Water hammer not well computed
- Pressurized sound speed $\simeq \sqrt{S/T_{\text{fente}}}$ \Rightarrow adjustment of T_{fente}
- Depression \Rightarrow seen as a free surface state

Preissmann (61), Cunge *et al.* (65), Baines *et al.* (92), Garcia-Navarro *et al.* (94), Capart *et al.* (97), Tseng (99)

OUR GOAL :

- Use Saint-Venant equations for free surface flows

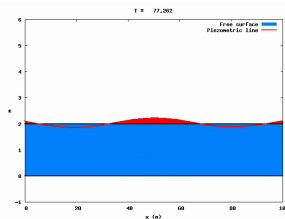
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- Use Saint-Venant equations for free surface flows
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 - ▶ which takes into account the depression
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- Get one model for mixed flows

To be able to simulate, for instance :



Ersoy et al. *A model for unsteady mixed flows in non uniform closed water pipes and a well-balanced finite volume scheme.*, Int. J. On Finite Volumes, 6(2), pp 1-47, 2009.

$$\begin{aligned}\rho_0 \operatorname{div}(\mathbf{U}) &= 0 \\ \rho_0 (\partial_t \mathbf{U} + \mathbf{U} \cdot \nabla \mathbf{U}) + \nabla p &= \rho_0 F\end{aligned}$$

Method :

- ① Write Euler equations in curvilinear coordinates.
- ② Write equations in non-dimensional form using the small parameter $\epsilon = H/L$ and takes $\epsilon = 0$.
- ③ Section averaging $\overline{U^2} \approx \overline{U} \overline{U}$ and $\overline{UV} \approx \overline{U} \overline{V}$.
- ④ Introduce $A_{sl}(t, x)$: wet area, $Q_{sl}(t, x)$ discharge given by :

$$A_{sl}(t, x) = \int_{\Omega(t, x)} dy dz, \quad Q_{sl}(t, x) = A_{sl}(t, x) u(t, x)$$

$$u(t, x) = \frac{1}{A_{sl}(t, x)} \int_{\Omega(t, x)} U(t, x) dy dz$$

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J.-F. Gerbeau, B. Perthame

Derivation of viscous Saint-Venant System for Laminar Shallow Water; Numerical Validation.
Discrete and Continuous Dynamical Systems, Ser. B, Vol. 1, Num. 1, 89–102, 2001.



F. Marche

Derivation of a new two-dimensional viscous shallow water model with varying topography, bottom friction and capillary effects.
European Journal of Mechanic B/Fluid, 26 (2007), 49–63.

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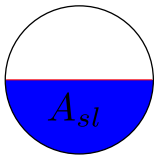
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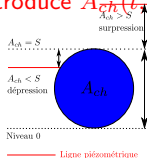
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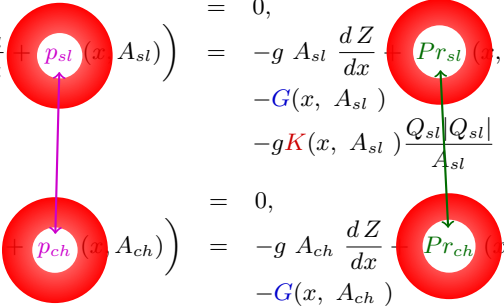
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Continuity criterion

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Continuity of **S** at transition point

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\longrightarrow

$$p(x, A, E) = c^2(A - \mathbf{S}) + gI_1(x, \mathbf{S}) \cos \theta$$

- Similar construction for the pressure source term :

$$Pr(x, A, E) = c^2 \left(\frac{A}{\mathbf{S}} - 1 \right) \frac{dS}{dx} + gI_2(x, \mathbf{S}) \cos \theta$$

$$\left\{ \begin{array}{l} \partial_t(A) + \partial_x(Q) \\ \partial_t(Q) + \partial_x \left(\frac{Q^2}{A} + p(x, A, E) \right) \\ \\ \\ \\ \end{array} \right. \begin{array}{l} = 0 \\ = -g A \frac{d}{dx} Z(x) \\ + Pr(x, A, E) \\ - G(x, A, E) \\ - g K(x, \mathbf{S}) \frac{Q|Q|}{A} \end{array}$$



Ersoy et al., *A model for unsteady mixed flows in non uniform closed water pipes.*, SCIENCE CHINA Mathematics, 55(1) :1–26, 2012.

MATHEMATICAL PROPERTIES

- The PFS system is **strictly hyperbolic** for $A(t, x) > 0$.
- For regular solutions, the mean speed $u = Q/A$ verifies

$$\partial_t u + \partial_x \left(\frac{u^2}{2} + c^2 \ln(A/S) + g \mathcal{H}(S) \cos \theta + g Z \right) = -g K(x, \mathbf{S}) u |u|$$

and **for** $u = 0$, we have :

$$c^2 \ln(A/\mathbf{S}) + g \mathcal{H}(\mathbf{S}) \cos \theta + g Z = cte$$

where $\mathcal{H}(\mathbf{S})$ is the physical water height.

- There exists a **mathematical entropy**

$$E(A, Q, S) = \frac{Q^2}{2A} + c^2 A \ln(A/\mathbf{S}) + c^2 S + g \bar{z}(x, \mathbf{S}) \cos \theta + g AZ$$

which satisfies

$$\partial_t E + \partial_x (E u + p(x, A, E) u) = -g A K(x, \mathbf{S}) u^2 |u| \leq 0$$

The PFS model have complexe source terms which makes its analysis hard !
 We have developped two numerical scheme :

- VFRoe (**Ersoy** et al., IJFV, 09)
- Kinetic (**Ersoy** et al., JSC, 11)

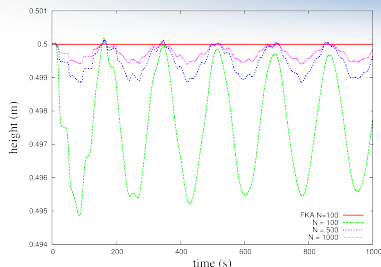
Nevertheless (for both), one can analytically check

- ① Positivity of A ,
- ② Conservativity of A ,
- ③ Drying and flooding phenomenon are naturally obtained
- ④ Discrete still water steady states are preserved exactly (only for VFRoe scheme)

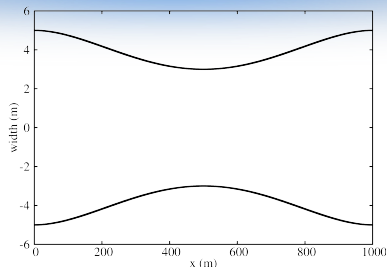
and numerically check that

- ① Discrete still water steady states are \approx preserved,
- ② Discrete in-entropy inequalities are *a priori* satisfied.

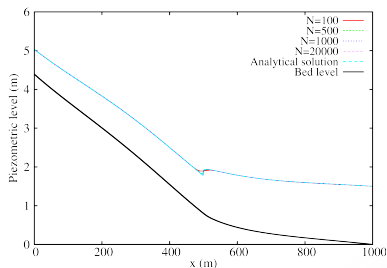
NUMERICAL ILLUSTRATION



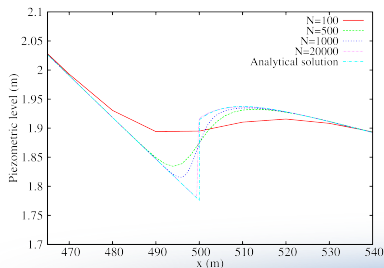
VFRoe : a still water steady states



Kinetic : variation of section



Kinetic : a transcritical steady solution with shock



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THE KINETIC FORMULATION

(A, Q) is solution of the PFS system if and only if \mathcal{M} satisfy the transport equation :

$$\partial_t \mathcal{M} + \xi \cdot \partial_x \mathcal{M} - g \Phi \partial_\xi \mathcal{M} = \mathcal{K}(t, x, \xi)$$

where $\mathcal{K}(t, x, \xi)$ is a collision kernel satisfying a.e. (t, x)

$$\int_{\mathbb{R}} \mathcal{K} d\xi = 0, \quad \int_{\mathbb{R}} \xi \mathcal{K} d\xi = 0$$

and Φ are the source terms.



B. Perthame.

Kinetic formulation of conservation laws.

Oxford University Press.

Oxford Lecture Series in Mathematics and its Applications, Vol 21, 2002.

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General form of the source terms :

$$\Phi = \overbrace{\frac{d}{dx} Z}^{\text{conservative}} + \overbrace{\mathbf{B} \cdot \frac{d}{dx} \mathbf{W}}^{\text{non conservative}} + \overbrace{K \frac{Q|Q|}{A^2}}^{\text{friction}}$$

with $\mathbf{W} = (Z, S, \cos \theta)$

- conservative term : classical upwind
- non conservative term : mid point rule (DLM, 95)
- friction : dynamic topography (Ersøy, Ph.D.)

1 INTRODUCTION

- Applications
- Derivation of SWE
- Properties of SWE and numerical illustration

2 NUMERICAL APPROXIMATION OF SWE

- Finite volume scheme for Homogenous SWE
- Finite volume scheme for non Homogenous SWE
- Kinetic scheme

3 SOME PHYSICAL APPLICATIONS DERIVED FROM SWE

- A sediment transport model
- An unsteady mixed flows in closed water pipes model (PFS model)

4 CONCLUDING REMARKS

Even if SWE are academic equations,

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- nevertheless, there are still several problems (even in the 1D case)
 - ▶ for conservative source terms (typically Z') : scheme preserving all steady states ?
 - ▶ for non conservative (as in the PFS model) : scheme preserving still water steady states exactly as well as an entropic scheme ?



M. Ersoy. *Modeling, mathematical and numerical analysis of various compressible or incompressible flows in thin layer – Modélisation, analyse mathématique et numérique de divers écoulements compressibles ou incompressibles en couche mince*, Ph.D., University of Chambéry, 2010.



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E. Toro, *Riemann solvers and numerical methods for fluid dynamics*, Springer-Verlag, 2009.



D. Serre, *Systems of conservation laws. 1. Hyperbolicity, entropies, shock waves.*, Cambridge University Press, 2010.

A dynamic background image showing a large splash of water with many droplets in the air, creating a sense of movement and freshness. The water is a clear, light blue color.

Thank you

Thank you

for your

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attention

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